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On the impact of the vertical resolution on chemistry transport modelling

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Abstract

This paper presents a sensitivity analysis of the modelling of air pollutant concentrations in the surface layer with the WRF/CHIMERE models. The influence of the vertical resolution near the surface is studied. The simulations are carried out over two periods (winter and summer 2009) over the Paris area. Three model configurations are used: (i) the CHIMERE mesh used for the PREVAIR forecast (8 levels from 995 to 500 hPa), (ii) a mesh refined along the whole vertical axis (20 levels from 995 to 500 hPa) and (iii) a mesh with a refinement near the surface (9 levels from 999 to 500 hPa). The results are discussed in terms of differences on surface concentrations between the reference case and an improved resolution. Adding a point close to the surface appears to be important mainly for high nocturnal concentrations in very stable boundary layers. Refining the vertical mesh, with 20 levels instead of 8, enables to model new structures in the well mixed boundary layer, but with a moderate impact at the surface. It is shown that the different model configurations lead to changes of a few $\mu\text{g}/\text{m}^3$ at most, showing that the vertical mesh is not the most sensitive factor in chemistry-transport modelling when results are compared to surface measurements. This finding validates the fact that a simplified vertical mesh is suitable for air quality forecasting even if an improved vertical resolution close to the ground is important to take into account the urban increment.

1. Introduction

Albeit the knowledge of pollution sources and meteorology are constantly improved, representing these processes in chemistry transport models (CTM) remains challenging. In particular, the representation of turbulence in the surface layer is difficult and has an important impact on the vertical mixing of meteorological parameters (Cionco and Ellefsen (1998), Kim et al. (2009)) and surface emissions (Pierce et al. (2010)). It influences the whole atmospheric chemical cycle and, in turn, the impact of pollution on human health (Imhof et al. (2005), Andersson et al. (2009), Valari et al. (2011)). In order to improve our knowledge of human exposure, it is necessary to improve the representation of pollutants in the boundary layer and more specifically within the first 100 meters above the ground (Lin and McElroy (2010)). This would give more realistic concentrations values where people live i.e. where the major part of primary pollutants are emitted and where air quality networks monitor pollutant concentrations, while models usually give a "mean vertical concentration value" around 20 to 50 m above ground level (agl).

There are various techniques used to retrieve more realistic values near the surface in the literature.

First, a parameterization can be developed to extrapolate 50 m concentrations (calculated by the CTM) to the surface (Byun and Dennis (1995)). In this case, the extrapolation has to take into account the difference between gaseous and particulate species, because of their chemical and physical characteristics and densities, leading

to different settling velocities and deposition rates (Xiao and Taylor (2002), Chamecki et al. (2007)). The challenge consists in building such a parameterization, knowing that pollutant concentrations are exposed to mixing and chemistry (Cheinet and Teixeira (2003)). There is thus no straightforward technique to extrapolate concentrations values down to the ground, because of the nonlinearities in the vertical profile of the concentration near the ground.

A second way would be to embed a Large Eddy Simulation (LES) or a local dispersion model in the CTM (Aristodemou et al. (2009), Garmory et al. (2009), Hara et al. (2009), Lee et al. (2008), Murena et al. (2009), Santos et al. (2009)). For example, the ADMS-urban (Righi et al. (2009)) model is sometimes used for urban pollution applications (Davies et al. (2007)). This is certainly the more realistic way in terms of fine turbulence modelling. This method is suitable for complex studies on restricted areas but the methodology is difficult to extend over large domains and not suitable for long term modelling or forecast purposes.

A third technique consists in adding discrete points in the vertical mesh at requested altitudes, or to refine this vertical mesh (Byun and Dennis (1995)). The advantage is that the chemical and meteorological behavior of the surface layer are maintained. This is crucial since most sources are emitted near the surface, such as nitrogen oxides emissions due to traffic in urbanised areas. The main goal of this paper is to estimate the impact of (i) increas-

ing the number of vertical model levels in the boundary layer, and (ii) changing the first model layer altitude, on pollutant concentrations.

In this study, three model configurations were defined and used to simulate two different periods over the Paris area (January and August 2009). The results are presented in terms of modelled concentrations differences between the configurations and discussed in terms of surface maps and vertical profiles for the usual regulated pollutants: ozone, nitrogen oxides and particulate matter (PM₁₀).

2. Models configurations

2.1. The meteorological model WRF

The regional model used in this study is the Weather Research and Forecasting (WRF) model in its version 3.2.1. This configuration was already used with CHIMERE for many studies such as Colette et al. (2011) and Menut et al. (2012b), among others. The model is used in its non-hydrostatic configuration. The coarse and fine horizontal domains include 90x70 and 58x58 grid points, respectively with 45km and 15km resolutions. The vertical domain covers 38 levels from the surface to 50 hPa and the integration time step is 40s for the fine domain. For the microphysics, the WRF Single Moment-5 class scheme is used allowing for mixed phase processes and super cooled water, (Hong et al. (2004)). The radiation scheme is RRTMG scheme with the MCICA method of random cloud overlap, (Mlawer et al. (1997)). The surface layer scheme is based on Monin-Obukhov with Carlson-Boland viscous sub-layer. The surface physics is calculated using the Noah Land Surface Model scheme with four soil temperature and moisture layers, (Chen and Dudhia (2001)). The planetary boundary layer physics is processed using the Yonsei University scheme, (Hong et al. (2006)) and the cumulus parameterization uses the Kain-Fritsch scheme for deep and shallow convection using a mass flux approach with downdrafts and CAPE removal time scale (Kain (2004)).

2.2. The chemistry-transport model CHIMERE

CHIMERE is an Eulerian off-line chemistry-transport model (CTM), as such it uses prescribed meteorological fields and emissions. The results presented in this paper were obtained with the 2011a version. Using these meteorological input data and a set of NO_x, SO_x, NH₃, PM, VOCs (Volatile Organic Compounds) and CO emissions, CHIMERE calculates the atmospheric concentrations of tens of gas-phase and aerosol species over local to continental domains (from 1 km to 1 degree resolution), and with a hourly output time step. All important chemical and physical processes are taken into account: emissions, transport (advection and mixing), chemistry and deposition. The concentrations fields are calculated for pollution event analysis, scenarios studies and forecasts (Menut and Bessagnet (2010)).

The dynamics and gas-phase mechanisms of the model are described in Schmidt et al. (2001), and improvements have successively been carried over (Vautard et al. (2005), Bessagnet et al. (2008)). The model documentation can be found on the web server, (www.lmd.polytechnique.fr/chimere). For both ozone and PM₁₀, the model has undergone extensive modelled aerosols intercomparisons at European and city scales (Vautard et al. (2007); van Loon et al. (2007); Schaap et al. (2007)). The aerosols model species are sulphates, nitrates, ammonium, secondary organic aerosols, primary particles and sea-salt (see Bessagnet et al. (2010) for details). The particle size distribution ranges from about 40 nm to 10 μm distributed in 8 bins. The gas-particle partitioning of the ensemble Sulfate/Nitrate/Ammonium is managed by the ISORROPIA scheme (Nenes et al. (1998)) implemented in CHIMERE.

The surface emissions account for both anthropogenic and biogenic emissions. The biogenic emissions are diagnosed using the MEGAN model (Guenther et al. (2006)). The anthropogenic emissions are based on the EMEP inventory and regridded over the specific CHIMERE domain, by reaggregating the chemical species following the gas-phase and particles chemical mechanism used (Menut et al. (2012a)).

2.3. Modelling domains

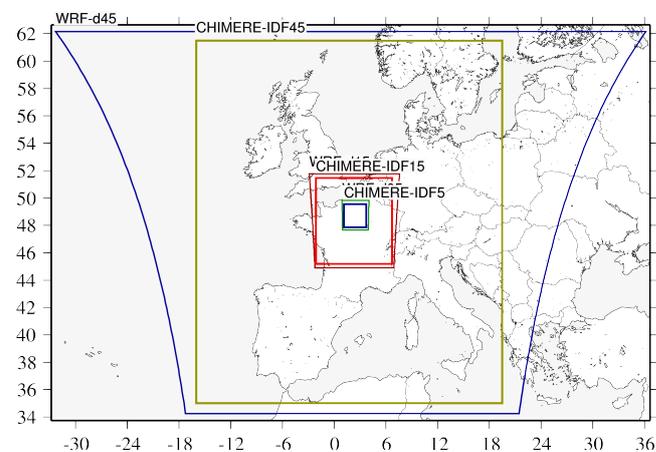


Figure 1: Model simulations domains used for WRF meteorological model and CHIMERE chemistry-transport model.

The simulation domains are centered on the Paris area, as shown in Figure 1. In order to calculate realistic pollutants concentrations, three different domains are nested and are described in Table 1. The largest domain, IDF45, has an horizontal resolution of $0.5^\circ \times 0.5^\circ$ and covers the whole western Europe. It is designed to describe long-range transport of pollutants and provides accurate boundary conditions to the two other domains. The IDF15 domain has a $15\text{km} \times 15\text{km}$ horizontal resolution, it is centered over Paris and covers the northern part of France.

The domain IDF5, has a $5\text{km} \times 5\text{km}$ horizontal resolution and is located at the center of IDF15. Figure 2 presents the land cover as used by CHIMERE. The city of Paris is considered as 100% urban. This percentage decreases sharply around the city and the suburb represents a circle around Paris with a radius of $\approx 20\text{km}$. For the rest of the domain, the region is mainly composed of agricultural lands, and, to a lesser extent, of grassland and forests.

In this paper, the results of the IDF5 domain are analysed, the IDF45 and IDF15 simulations provide exclusively meteorological and chemical boundary conditions to the IDF5 domain.

Domain	$N_x \times N_y$	$\Delta x \times \Delta y$
<i>IDF45</i>		
WRF	72×54	$45\text{km} \times 45\text{km}$
CHIMERE	68×50	$0.5^\circ \times 0.5^\circ$
<i>IDF15</i>		
WRF	49×52	$15\text{km} \times 15\text{km}$
CHIMERE	45×48	$0.2^\circ \times 0.133^\circ$
<i>IDF5</i>		
WRF	46×49	$5\text{km} \times 5\text{km}$
CHIMERE	38×35	$0.07^\circ \times 0.05^\circ$

Table 1: Model domains set-up as displayed in Figure 1. The WRF model uses a Lambert projection and CHIMERE a Lambert conformal projection

2.4. Landuse

There are currently 9 land use categories in CHIMERE. Those categories are calculated from available global land use databases, which can contain different number of classes. In this version, we used the GlobCover Land Cover database, a global land cover map at 10 arc second (300 meter) resolution (Bicheron et al. (2008)). It contains 22 global land cover classes defined within the UN Land Cover Classification System (LCCS). The GlobCover database is based on the ENVISAT satellite mission’s MERIS sensor (Medium Resolution Image Spectrometer) Level 1B data retrieved in Full Resolution (FR) mode with a spatial resolution of 300 meters. GlobCover LC was derived from an automatic and regionally-tuned classification of a time series of MERIS FR composites covering the period December 2004-June 2006.

Figure 2 [top] presents the domain landuse in each cell and over the Paris area. Over the whole domain, the ‘Agricultural land / crops’ is the main landuse type, when the center of the domain is fully urbanised with the Paris city.

2.5. Vertical meshes

Three vertical meshes are defined for the sensitivity studies performed with the CHIMERE model. The altitudes of the vertical levels are displayed in Table 2. The three model configurations are described hereafter:

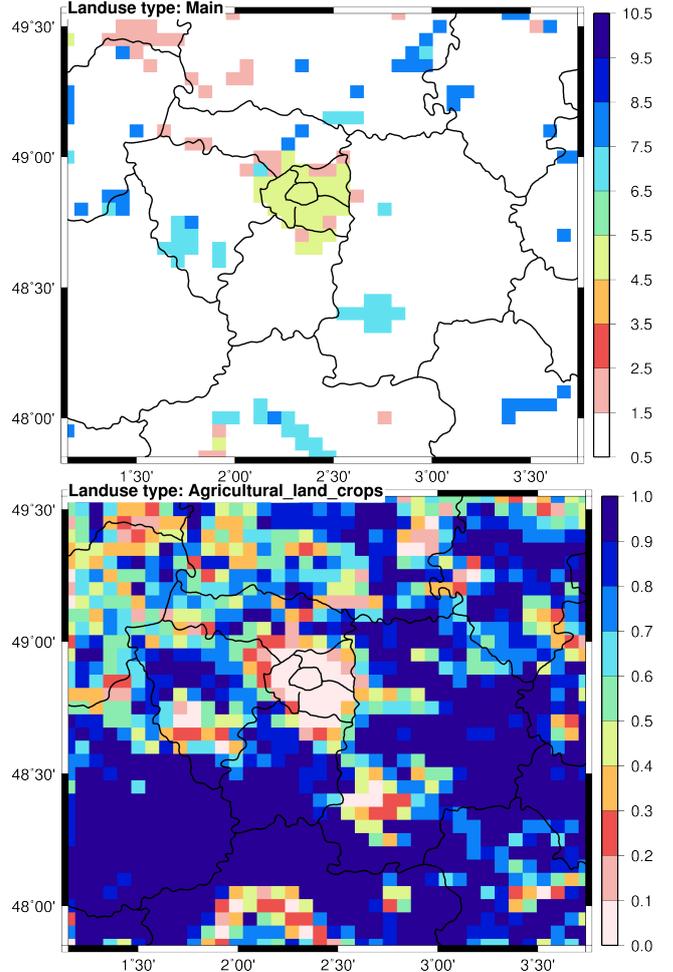


Figure 2: [top] Landuse map for the Paris region: for each cell the dominant landuse is drawn. The color code corresponds to: 1 Agricultural land / crops, 2. grassland, 3. barren land, 4. inland water, 5. urban, 6 shrubs, 7 needleleaf forest, 8 broadleaf forest, 9 ocean. [bottom] the relative percentage [0:1] of the first sector, Agricultural land and crops.

- **C8**: Eight vertical levels are defined from 995 to 500 hPa, representing the lower troposphere, and starting with the first vertical point at about 40m above ground level. This corresponds to the operational forecast configuration designed for the PREVAIR system (Rouil et al. (2009)). This configuration is considered in this study as the reference case.
- **C20**: Twenty vertical levels are defined, also from 995 to 500 hPa. This version is designed to examine the impact of a higher resolution mesh compared to C8.
- **C9**: Nine vertical levels are defined. This is strictly the same as C8 but adding one point near the surface located at 999 hPa. This configuration aims at quantifying the variability in the surface concentrations results by adding a point near the surface which would be more representative of measurements.

Label	Altitude of levels tops (m)
WRF30	27; 93; 153; 298; 443; 627; 852; 1157; 1531; 1918; 2320; 2910; 3690; 4485; 5273; 6050; 6850; 7640; 8420; 9200; 10010; 11700; 12600; 13480; 14420; 15370; 16400; 17400; 18500; 19700
C8	40; 108; 226; 431; 790; 1438; 2654; 5191
C9	8; 40; 108; 226; 431; 790; 1438; 2654; 5191
C20	40; 86; 139; 200; 271; 353; 448; 559; 687; 836; 1012; 1216; 1457; 1741; 2075; 2479; 2967; 3558; 4278; 5191

Table 2: Top layers altitudes in meters for the WRF meteorological model (30 levels) and for the three CHIMERE configurations: C8 (8 levels), C9 (9 levels) and C20 (20 levels).

The results will be presented in terms of differences in modelled concentrations fields compared to surface measurements. The following behaviours are expected for the three configurations:

- The C8 and C20 vertical meshes have the same first level point. The only difference is that C20 has more discrete values in the boundary layer (the number of points in the boundary layer is not constant since it depends on the boundary layer height). By difference between C8 and C20 concentrations fields, the results will quantify the impact of a vertical refinement and thus the impact of the accuracy of meteorological vertical gradients within the boundary layer.
- The C8 and C9 vertical meshes have the same vertical mesh, but C9 has an additional point near the surface. In this case, the mixing and transport within the surface layer would have the same behaviour, but with C9 the model will calculate explicitly a more realistic value at the ground level.

In the next sections, results will be presented either as absolute values for each of the three configurations, or as differences: in that case the C8 configuration is considered as the reference because it is used in operational forecast mode in the PREVAIR system, Rouil et al. (2009), Honoré et al. (2008), Menut and Bessagnet (2010).

2.6. Estimation of turbulent parameters

CHIMERE offers the possibility to use either diagnostic available variables from meteorological drivers or diagnose some meteorological variables in the pre-processing step. In this study, the latter option is selected in order to use the same turbulent parameterizations and also to be consistent with the CHIMERE model configuration used in the PREVAIR system (Rouil et al. (2009)). The full WRF vertical resolution is used to diagnose the vertical turbulent diffusivity K_z and the boundary layer height \bar{h} . After this evaluation, the diagnosed profiles of K_z are averaged over the CHIMERE vertical grid in the same way as all others meteorological parameters. It is during this final

vertical averaging that the number of vertical levels used by the CHIMERE model are changed.

The following variables are calculated: the friction velocity u_* , the surface sensible heat flux Q_0 , the vertical convective velocity w_* , the boundary layer height \bar{h} , the Monin-Obukhov length L and the vertical diffusivity profile K_z .

The friction velocity u_* is used for deposition and the calculation of diffusivities. It is a particularly sensitive parameter for ozone in summer through the calculation of aerodynamic resistance r_a . It depends strongly on land use types which are critical to deposition. In large scale meteorological models, roughness lengths are sometimes too coarse and cannot be applied for high-resolution deposition. Recomputing u_* as proposed by Louis et al. (1982) gives satisfactory results. In addition it yields a deposition that is consistent with the high-resolution land use.

$$u_* = \sqrt{C_{DN}^2 F_m |u(z)|^2} \quad (1)$$

where F_m is the Louis et al. (1982) stability function, z the height above ground level where the meteorological data are available (in our case $z=10\text{m}$). C_{DN} the neutral drag coefficient as:

$$C_{DN} = \frac{k}{\ln\left(\frac{z + z_{0m}}{z_{0m}}\right)} \quad (2)$$

with $k=0.41$, the Karman constant. z_{0m} is the dynamical roughness length (in meters) and values are tabulated for each landuse type. F_m is the momentum stability function and is estimated depending on the bulk Richardson number value. Under stable cases (if $R_{ib} < 0$):

$$F_m = 1 - \frac{2bR_{ib}}{1 + 3bcC_{DN}^2 \sqrt{\frac{z + z_{0m}}{z_{0m}}} \sqrt{|R_{ib}|}} \quad (3)$$

and under unstable cases (if $R_{ib} > 0$):

$$F_m = \frac{1}{1 + \frac{2bR_{ib}}{\sqrt{1 + dR_{ib}}}} \quad (4)$$

with the constant $b = c = d = 5$. Under neutral case, i.e $R_{ib} = 0$, $F_m=1$.

The bulk Richardson number is estimated as:

$$R_{ib}(z) = \frac{g z}{\theta_v(z)} \frac{\Delta \bar{\theta}_v}{\Delta |U|^2} \quad (5)$$

with $\Delta \bar{\theta}_v = \bar{\theta}_v(z) - \bar{\theta}_v(z_0)$, $\Delta |U| = |U|(z)$.

Heat fluxes are used for the calculation of w_* (and therefore the vertical mixing), and the height of the boundary layer. Only the virtual heat flux is required, which can be recalculated from an empirical formula Priestley (1949) using temperatures in the first meteorological model layers. Because of uncertainties in the later formula, it is strongly

advised to use the heat fluxes of the meteorological model if available. If the surface sensible heat fluxes Q_0 are provided by the meteorological model, they are directly used for the calculation of the convective velocity w_* :

$$w_* = \left(\frac{gQ_0\bar{h}}{\rho C_p \bar{\theta}_v} \right)^{1/3} \quad (6)$$

where $g=9.81 \text{ m}^2.\text{s}^{-2}$, Q_0 is the surface sensible heat flux, \bar{h} is the convective boundary layer height, C_p is the specific heat of air at constant pressure, $\bar{\theta}_v$ is the mean virtual potential temperature representative of the surface layer.

The Monin-Obhukov length is estimated as:

$$L = \frac{-\bar{\theta}_v u_*^3}{k g Q_0} \quad (7)$$

The boundary layer height (\bar{h}) differs for stable and unstable conditions. Under stable conditions, i.e when $L > 0$, \bar{h} is estimated as the altitude where the Richardson number reaches a critical number here chosen as $R_{ic} = 0.5$, following Troen and Mahrt (1986).

Under unstable (i.e convective) conditions, \bar{h} is also chosen as the altitude where $R_i = R_{ic}$, including the influence of boundary layer convective clouds. The latter is based on a simplified and diagnostic version of the approach of Cheinet and Teixeira (2003). It consists in the resolution of the (dry) thermal plume equation with diffusion. The in-plume vertical velocity and buoyancy equations are solved and the boundary layer is taken as the height where the vertical velocity stops. Thermals are initiated with a non-vanishing vertical velocity and potential temperature departure, depending on the turbulence similarity parameters in the surface layer.

A usual approximation for CTMs, considered as numerically diffusive, is to neglect the horizontal turbulent fluxes. Vertical turbulent mixing takes place only in the boundary-layer. The formulation uses K-diffusion following the parameterization of Troen and Mahrt (1986), with no counter-gradient term. In each model column, the diffusivity K_z (m^2/s) is calculated as:

$$K_z = k w_s \frac{z}{\bar{h}} \left(1 - \frac{z}{\bar{h}} \right)^2 \quad (8)$$

In this equation, w_s is a vertical scale given by similar formulas:

- In the stable case (when the surface sensible heat flux is negative):

$$w_s = \frac{u_*}{(1 + 4.7z/L)} \quad (9)$$

- In the unstable case:

$$w_s = (u_*^3 + 2.8 e w_*^3)^{1/3} \quad (10)$$

where $e = \max(0.1, z/\bar{h})$. A minimal $K(z)$ is assumed, with a value of $0.01 \text{ m}^2/\text{s}$ in the dry boundary layer and of $1 \text{ m}^2/\text{s}$ in the cloudy boundary layer. In the boundary layer, a maximal value of $K_z=500 \text{ m}^2/\text{s}$ is used to avoid unrealistic mixing. Above the boundary layer, a fixed value of $K_z=0.1 \text{ m}^2/\text{s}$ is prescribed.

3. Overview of the two periods

3.1. Meteorological conditions

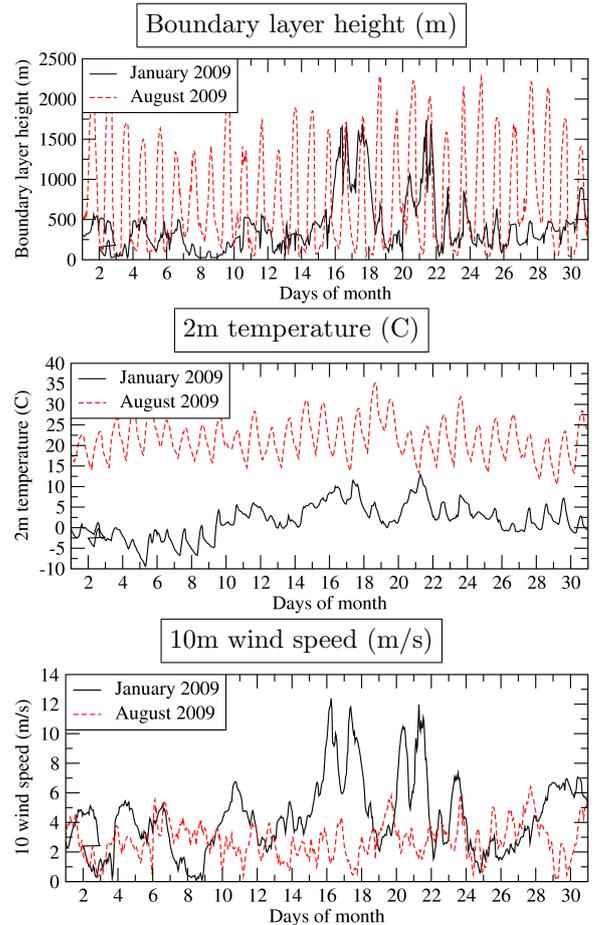


Figure 3: Times series of the mean meteorological parameters for the period of January and August 2009 at the center of the domain: Paris.

The results in this study are presented for the 1-31 January and 1-31 August 2009 periods. These periods consist of one winter and one summer month over the Paris area, in order to cover several meteorological situations leading to pollution events in this region. These periods were selected for the following reasons: (i) they correspond to "classical" meteorological situations leading to normal pollution episodes; (ii) no precipitations were recorded, leading to a more robust analysis of the modelled vertical mixing of pollutants, (iii) in the two cases, pollution events were reported by the local air quality networks.

Time series of meteorological parameters such as 2m temperature, boundary layer height and 10m wind speed are presented in [Figure 3](#). These parameters correspond to the meteorological parameters used by the CHIMERE model. They are fully independent from the vertical resolution chosen for the chemistry-transport modelling. The boundary layer heights are very different between January and August: the diurnal cycle is rather low during January when values quickly evolve during the daytime in August. In January, the mean values are $\approx 400\text{m}$, and maximal values of $\approx 1500\text{m}$ are observed between the 15 and 22 January. In August, the boundary layer top is always between 1500 and 2500m. The 2m temperature also exhibits low diurnal cycles during January and a high daily amplitude in August. The mean values are ≈ 0 and 23°C for January and August, respectively. In summer, the nocturnal values are relatively low (from 10 to 15°C). The 10m wind speed is low in August, with a mean value of $\approx 3\text{m/s}$ favourable to stagnation and, thus, photochemical pollution episodes. During winter, the 10m wind speed may rapidly change, from very low wind periods (e.g. 8 January) to high surface wind speed days (e.g. 16, 18, 22 January). These high wind speed values of mid-January explain the high boundary layer height diagnosed for that period.

3.2. Pollutant concentrations

The first model level concentrations maps are displayed in [Figure 4](#) for NO_2 , PM_{10} and O_3 (in $\mu\text{g}/\text{m}^3$). The model configuration considered as the reference case is the C8 simulation ([Honoré et al. \(2008\)](#)). The maps represent monthly average concentrations over January and August 2009. Even if absolute values are different between summer and winter, the regional patterns are similar for the three pollutants O_3 , NO_2 and PM_{10} . The Paris area is at the center of the domain and exhibits the major source of primary pollutants such as PM_{10} and NO_2 . The limited spread around the city correspond to the mean concentrations transported a few kilometers out of the city center, and each pattern represents the mean wind direction during each month.

For both periods, the maximum values of NO_2 and PM_{10} are observed over Paris, characterised by the most important anthropogenic emissions fluxes. Due to fast chemical reactions (mainly titration), Paris is also the location of the lowest surface ozone values: $\approx 20 \mu\text{g}/\text{m}^3$ in January and $\approx 70 \mu\text{g}/\text{m}^3$ in August.

Around Paris, background values are homogeneous and lower, averaging the potential ozone plumes and the background concentrations (local biogenic production or long-range transport, [Menut et al. \(2000\)](#), [Derognat et al. \(2003\)](#)). The main differences are observed for ozone, the background averaged values are $\approx 40\text{-}50 \mu\text{g}/\text{m}^3$ in winter and $\approx 70\text{-}80 \mu\text{g}/\text{m}^3$ in summer.

4. Impact of a higher vertical resolution

4.1. Monthly averaged surface concentrations

The comparisons are presented for the differences (C20-C8) and (C9-C8), for the three studied pollutants, O_3 , NO_2 and PM_{10} , and for January and August 2009. Results are displayed in [Figure 5](#) and for the first vertical level only: it represents the "surface" concentrations i.e the mean value between the ground and the top layer altitude. Such concentrations are the values usually compared to the air quality measurements stations. For (C20-C8), the model configurations have the same first vertical level but not the same vertical resolution. A direct comparison between the first model level horizontal fields gives the direct impact of an increased vertical resolution. For (C9-C8), the first vertical levels are not at the same altitude above ground level: 8 and 40m respectively. It shows the potential improvement of adding a lower first level close to the observations (i.e usually $\approx 3\text{m}$).

Globally, the maps represent surface concentrations and not surface emissions. The most important emissions are located in the Paris city center (the center of the modelled domain). Due to horizontal transport and chemistry, the main differences are observed (i) in the south-west of Paris during January and August 2009, due the meteorology during this period. Thus, the most important impact appears to be in the suburb but has to be directly rely to the changes in the Paris city center.

For NO_2 surface concentrations ([Figure 5](#), left), the differences are very low and never exceed $\pm 3 \mu\text{g}/\text{m}^3$. In January, the main negative differences (i.e. C8 > C20) are observed in the center of the Paris area, while the main differences of positive sign are observed at the border of the domain. For the same period, the contrary is observed for (C9-C8), more logically directly linked to the fact that the compared vertical levels are different and that the NO_2 concentrations are always more important near the surface due to traffic emissions. In August, the (C20-C8) values are close to zero due to the photochemistry for the same vertical level and close to the ground. But some non negligible values may be observed with (C9-C8), with higher values for C9, for the same reason than during January.

For PM_{10} surface concentrations, the differences are larger and C20 shows concentrations higher by $+0.5$ to $+2 \mu\text{g}/\text{m}^3$ all over the domain in August and January, respectively. The differences are spatially more homogeneous over the domain for PM_{10} than for NO_2 . The main components of PM_{10} are a primary anthropogenic fraction and a secondary biogenic origin.

For O_3 concentrations, the differences are always negative, so that C8 > C20 over the whole domain. For January and August, the differences range is -3 to $-1 \mu\text{g}/\text{m}^3$. The differences are not spatially homogeneous and have to be directly related to the spatial patterns observed for NO_2 over the main source of the Paris city. The O_3 deposition in the C9 simulations is enhanced over forested

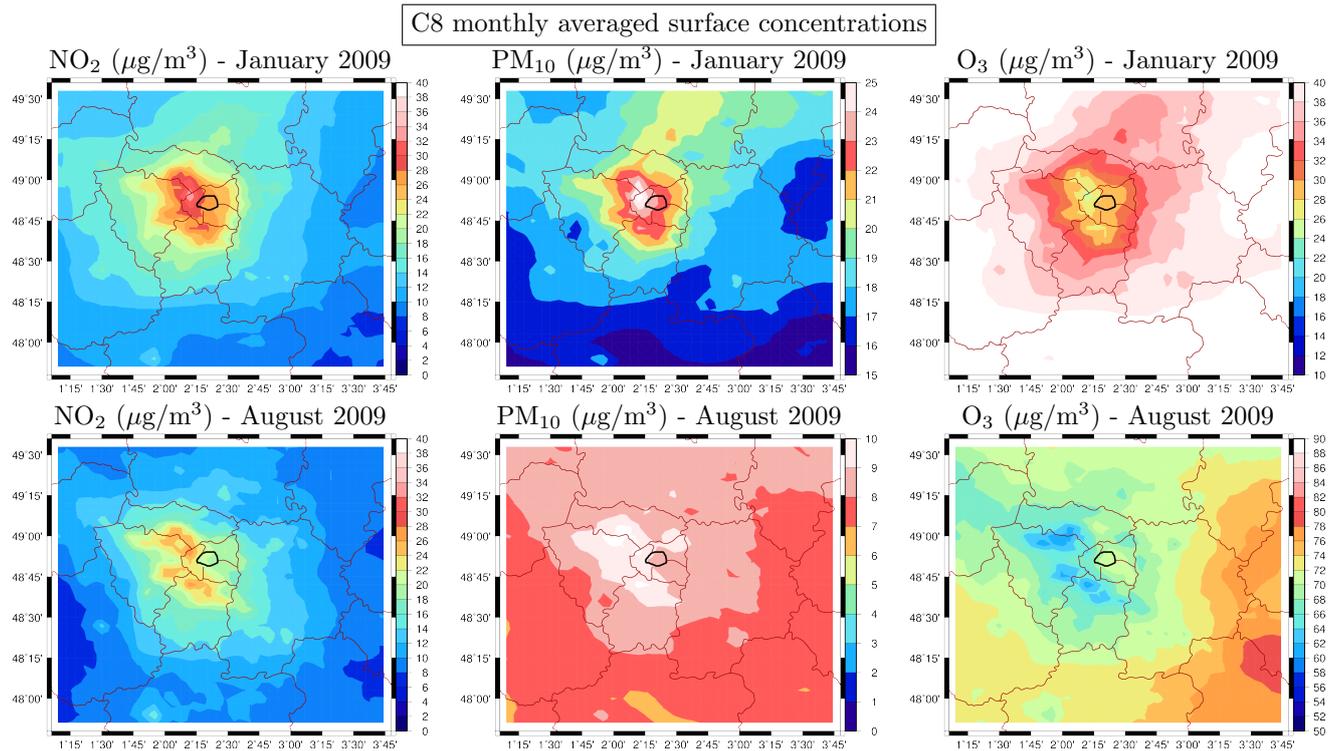


Figure 4: Maps of monthly mean averaged surface concentrations for NO_2 , PM_{10} and O_3 , and for the months of January and August 2009.

areas, lower concentrations are simulated compare to the reference simulation C8.

For the three pollutants, the differences reach on average $\pm 3 \mu\text{g}/\text{m}^3$. For PM_{10} and O_3 , Figure 5 clearly shows that refining the whole vertical mesh has a large scale impact, while adding one level near the surface has a more variable impact at the local scale depending on emissions and landuses.

4.2. Vertical cross sections

In this section, the analysis focussed on the vertical structure. Vertical time series are displayed to estimate the impact of the different vertical meshes on the whole troposphere concentrations. Two examples are chosen: the vertical profiles of PM_{10} during winter (January 2009, Figure 6, left panel) and ozone during summer (August 2009, Figure 6, right panel). As for all following results, the modelled concentrations are extracted from the model cell the most affected by the anthropogenic emissions, the Paris city center (longitude 2.35°W and latitude 48.86°N).

The comparison is purposely limited to a qualitative discussion. Since the vertical resolution is different, a quantification would require to interpolate the pollutant concentrations at identical altitudes. This interpolation process at such low native resolution would generate more differences than the differences observed between the three models configurations. Therefore, the results are thus just qualitatively discussed in this section.

The vertical PM_{10} profiles showed that the largest concentrations are simulated close to the surface, due to anthropogenic emissions. During the month of January 2009, the most important concentrations are observed during the first two and the last weeks. The lowest concentrations are observed during the period of 14 to 22 January 2009, being those when the highest boundary layer height and 10m wind speed values are modelled (Figure 3): surface concentrations are vertically mixed and/or advected out of the modelled domain. The impact of C20 appears clearly around the 20th of January, when a concentration of $\approx 20 \mu\text{g}/\text{m}^3$ is modelled between 800 and 1000m, but not with C8 and C9. The same type of structure, only modelled by C20, is also observed on January, 27 between 900 and 1000m.

The vertical profiles of ozone are different, ozone has always larger concentrations in the upper boundary layer than near the surface, when fast titration by NO_x occurs. The C20 configuration shows that the vertical gradients are more pronounced than for C8 and C9, mainly around 3000m, at the top of the boundary layer. But, on average, the most important vertical structures are modelled and the number of vertical levels does not seem to impact much the behaviour of ozone in the boundary layer. The most important differences are observed up to 4000m, showing that a finer resolution has an impact on the representation of long-range transport.

The comparison of vertical profiles for the three configurations showed that C20 is able to model ozone and

PM₁₀ in the boundary layer with differences up to 10 to 20 $\mu\text{g}/\text{m}^3$. These differences are often above the boundary layer and are related to long-range transport rather than local production. This feature shows that the vertical refinement has a moderate impact on surface concentrations.

4.3. Analysis of monthly mean vertical profiles

In order to better evaluate the impact of the vertical resolution on concentrations profiles, the concentrations displayed in Figure 6 are averaged on a monthly basis and presented for each model configuration in Figure 7. The upper panel presents K_z vertical profiles for two hours, 06:00 UTC and 15:00 UTC. For 06:00 UTC, for example, this means that all profiles of January and August at 06:00 UTC are averaged to obtain one profile for each month, representative of this specific hour. For O₃ and NO₂ pollutants, all hourly profiles of the month are averaged.

The K_z profiles show different shapes and magnitudes between 06:00 and 15:00 UTC. At 06:00 UTC, before the convective period, the values of diffusivity are low and do not exceed 30 and 70 m^2/s , at $\approx 300\text{m}$ and 1000m , in January and August, respectively. The shape of the profile is noisy because of the day to day variability. Note that the standard deviation was not found as a relevant information here, its values being of the same order of magnitude as the mean profiles. The K_z values are more important during summer than the winter, and may reach up to 160 m^2/s in August at 15:00 UTC, for a mean altitude of 1000m. For all times and periods, the most important differences are in the lowest layer in the C9 simulation, showing that the extra point close to the surface generates a very low mixing ($K_z < 2 \text{m}^2/\text{s}$), whilst the two other simulations always display minimum values between 5 and 10 m^2/s .

The NO₂ profiles show maximum values close to the surface, where emissions take place. The differences between C8 and C20 are rather low, as previously showed with the horizontal maps in Figure 5. The differences between C9 and the two other configurations are important, the concentrations gradually increase close to the ground. This leads to an increase of NO₂ concentrations of $\approx 1 \mu\text{g}/\text{m}^3$ in January and $\approx 3 \mu\text{g}/\text{m}^3$ August 2009.

For ozone, concentrations usually increase with altitude. This leads to lower values near the surface with C9. The differences are very low and never exceed a few $\mu\text{g}/\text{m}^3$ compared to C8 and C20. Up to 40m, the C8 and C9 profiles are similar and show large differences compared to the C20 profile. The C20 concentrations are lower between 40m and 1500m, but higher up to 1500m. The same impact is observed on the January profile, but with an inversion of differences around 700m. This feature reflects the impact of a refined vertical transport, between the boundary layer and the free troposphere.

The monthly mean concentrations profiles showed that the C8 and C9 configurations have the same values down to 40m. The additional point with C9, around 8m above

ground level, appears as an extrapolation of the upper profile. This leads to higher primary pollutants (here NO₂) and lower ozone concentrations at the first model vertical level. However, differences remain low and an additional extrapolation to 1 or 2m above ground level would not give large variations in surface time series of pollutants.

4.4. Sensitivity to the vertical mixing

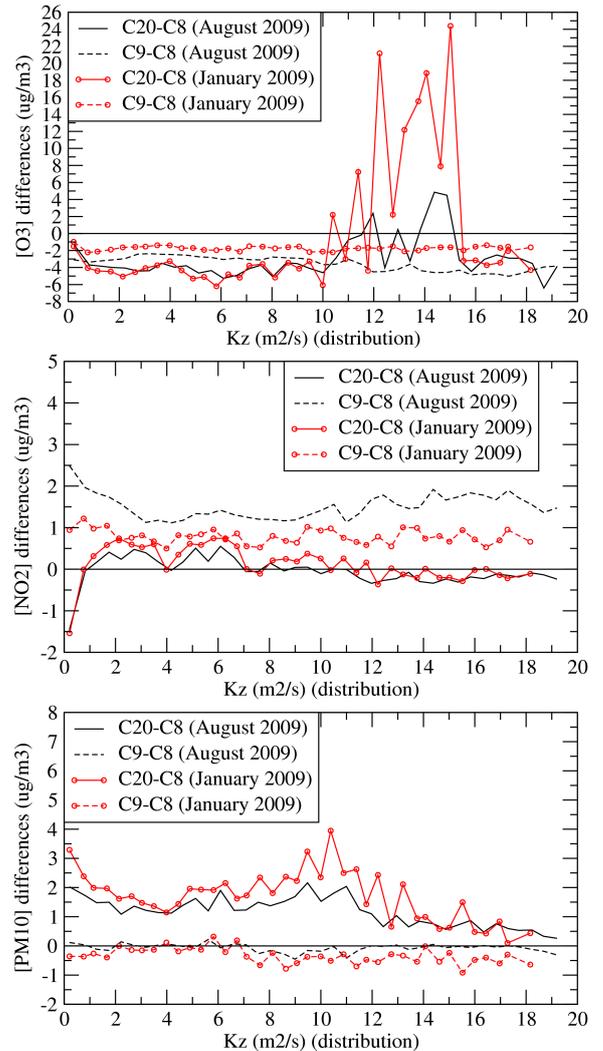


Figure 8: Distributions of concentrations differences as a function of the K_z first model level values. The K_z profiles being different between the model configurations, the C20 configuration was used here for the distributions calculations.

In order to quantify the impact of K_z on surface concentrations, the simulations are compared by ranking the concentrations differences, (C20-C8) and (C9-C8), as a function of the corresponding value of surface K_z (during August 2009). The surface K_z values are discretised into regular bins of 1 m^2/s step. For each K_z bin, the corresponding (C20-C8) and (C9-C8) hourly concentrations are averaged. The result gives an indication about the

surface K_z value for which the most important concentrations differences are obtained. The results are displayed in Figure 8, for O_3 , NO_2 and PM_{10} .

For ozone concentrations, the (C20-C8) and (C9-C8) differences for the two months are always negative in low K_z conditions. The differences do not exceed $-5 \mu\text{g}/\text{m}^3$. This means that in case of low mixing, adding a point near the surface or increasing the vertical resolution tends to have lower surface ozone concentrations. For higher K_z values, between 10 and $16 \text{ m}^2/\text{s}$, the differences turn positive and larger (up to $25 \mu\text{g}/\text{m}^3$ for the (C20-C8) and during January 2009. In the case of ozone, when the vertical diffusivity is high, during convective conditions, there is a competition between an enhanced titration close to the surface and an import of ozone coming from the upper layers. The large variability observed for ozone is not diagnosed for NO_2 and PM_{10} . The distributions are relatively flat and tend to show that the differences between the model configurations are not really sensitive to the mixing values near the surface.

For NO_2 , the values of differences are mainly positive, on average between 0 and $+2 \mu\text{g}/\text{m}^3$, showing that C9 and C20 tends to increase NO_2 concentrations compared to the reference run C8. Negative values are only observed for very low mixing ($K_z < 1 \text{ m}^2/\text{s}$) and for the C20 configuration. In this case, the refined vertical meshes gives lower NO_2 , but with moderate differences (with a maximum of $-1.5 \mu\text{g}/\text{m}^3$).

For PM_{10} , the differences are always close to zero between C8 and C9. For this primary/secondary pollutant, whatever the K_z value, i.e the stability conditions, there is a competition effect between chemistry and dynamics, that explains this feature. The chemistry leads to negative differences (as for ozone) and the dynamics to negative differences (as for NO_2). The differences are higher between C8 and C20, with values up to $+3 \mu\text{g}/\text{m}^3$, showing that an increase of the vertical resolution will increase surface PM_{10} concentrations.

5. Model to data comparison

In the previous sections, the impact of a change in the vertical resolution was discussed by comparing various model configurations. Here we evaluate whether these changes give better modelled results when compared to measurements. These comparisons are performed thanks to the measurements issued from the local air quality network operating over the Paris area (AIRPARIF). Hourly measurements are used to evaluate the day to day variability and the change of diurnal cycles in the various model configurations.

5.1. Surface time series

Many previous studies have been carried out regarding the comparison of modelled surface concentrations with CHIMERE in the Paris area: the ESQUIF experiment

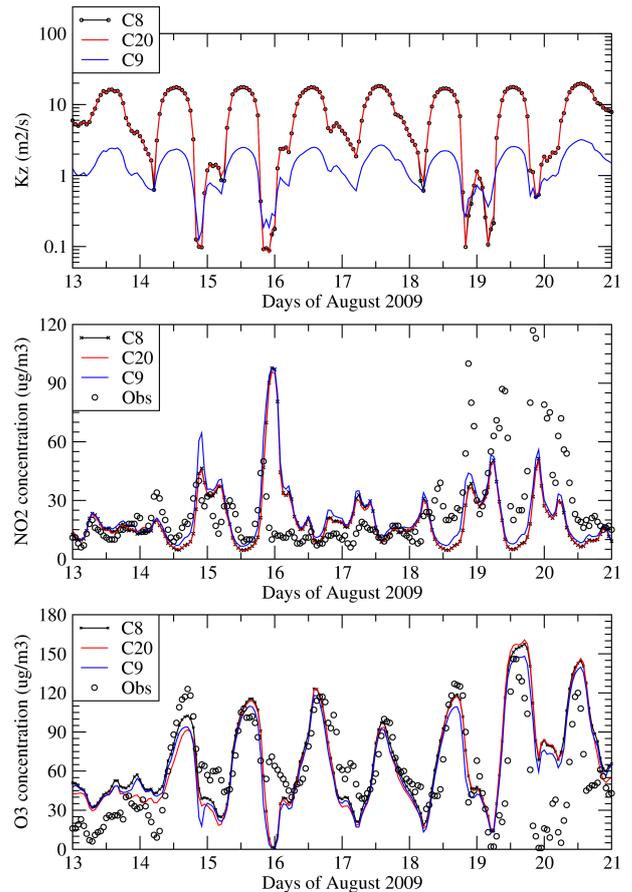


Figure 9: Example of times series of vertical turbulent diffusivity, ozone, O_3 , and nitrogen dioxide, NO_2 , at the first vertical level and during August 2009. Measurements and modelled concentrations represent the Paris center.

Menut et al. (2000), through the routine validation of the PREVAIR forecast system Rouil et al. (2009) or the sensitivity study of the horizontal resolution Valari and Menut (2008). It was shown that CHIMERE can accurately simulate and forecast ozone and NO_2 concentrations. Since the goal of this paper is not to discuss again the model validation but to evaluate its sensitivity to the vertical resolution, this section does not contain a full quantitative performance analysis but more a qualitative discussion of a change in the horizontal resolution and K_z values on modelled pollutant concentrations.

Figure 9 displays time series at the first vertical model level of the turbulent diffusivity coefficient, K_z , ozone and nitrogen dioxides concentrations. In order to focus on diurnal cycle changes, time series are restricted to the time period between the 13 and 21 August 2009, representative of a large day to day variability. The comparisons are done between the Paris grid cell for the model and the AIRPARIF network measurements, for the NO_2 et O_3 measurements data.

Figure 9 (top) shows the time series of K_z for the three configurations, C8, C20 and C9 and with a log scale along

the y axis in order to have a focus on low values. Having the same first vertical level, the values for C8 and C20 are exactly the same: the diurnal variability ranges from 0.1 to 2 m²/s during the night (stable conditions) to 15-20 m²/s during the afternoon (fully convective conditions). For C9, the same kind of diurnal variability is observed but with lower values, between 0.1 and 4 m²/s. For all configurations, very low K_z values are diagnosed for the nights of 14-15, 15-16 and 19-20 August 2009. The minimum K_z of 0.1 m²/s is reached for C8 and C20, but not for C9 (with higher values, 0.2 m²/s for the same time).

The comparison between modelled and measured NO₂ concentrations (Figure 9 (middle)) shows that the period is well modelled except (i) for the night of 15-16 August when the model overestimates the concentrations and, the period from 19 to 20 August, when the model underestimates the concentrations. In general, the most important NO₂ surface concentrations are modelled during the night and the morning when the K_z are the lowest. But, the fact that some K_z were very low in the model has no a systematic impact on surface concentrations: for the night of 15-16 August, the NO₂ overestimation can be attributed to the very low corresponding K_z values. This is not the case for the night of 19-20 August when the model underestimates the NO₂ while K_z values are also very low.

For ozone (Figure 9 (bottom)), the day to day variability is well captured by the model, and the major daily peaks are well captured. The main differences between the three configurations occur during the night, when the K_z differ. This is the case of the 15-16 August night, when the modelled ozone concentrations tends to zero while the observations are about 50 μg/m³. This is a direct impact of the low K_z values, producing a low mixing and thus overestimating NO₂ concentrations and O₃ titration. For night of 19-20 August, the underestimation of NO₂ concentration leads to non-zero ozone concentrations while the observations are close to zero.

The analysis of surface concentrations time series shows that the main impact of the vertical resolution occurs at night, when K_z is very low and similar to the minimum value imposed in the model. These low values may induce very large overestimation of primary species such as NO₂, and in turn an underestimation of nocturnal ozone.

Even if the change of the vertical resolution involves differences in modelled surface concentrations, these differences remain low compared to the differences between model and measurements. It appears obvious that the change in vertical resolution or the height of the first vertical level is not the most sensitive parameter in the modelling system even if refining the vertical mesh has a significant beneficial impact.

5.2. Diurnal cycles and comparisons to observations

In order to have an overview of the main differences, diurnal cycles are calculated for the whole months of January and August 2009. Cycles are calculated for ozone,

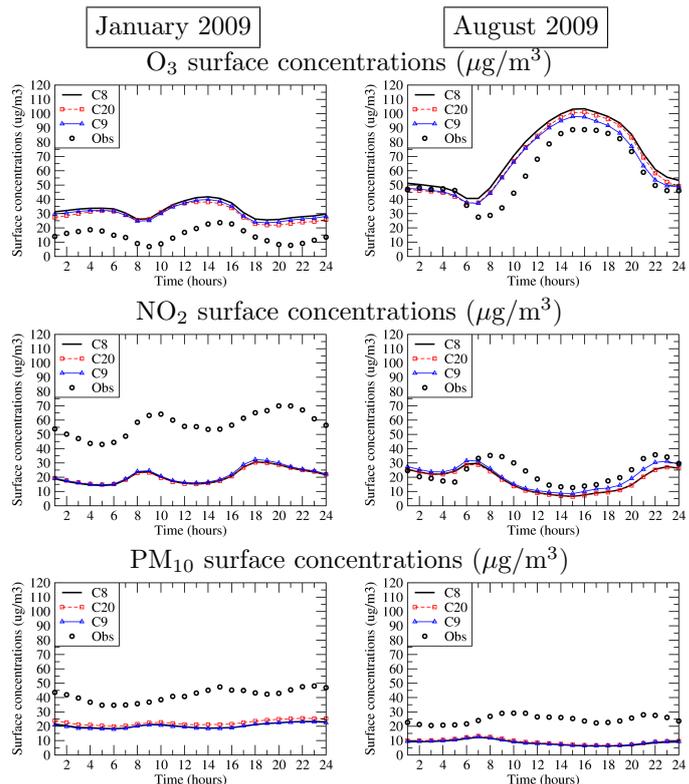


Figure 10: Diurnal cycles for surface concentrations of ozone, NO₂ and PM₁₀, January and August 2009 for Paris.

NO₂ and PM₁₀ for the three model configurations and for the corresponding surface measurements in the center of Paris. The same cycles were also calculated for other locations in the Paris area and similar behaviors were observed: the discussion is thus focusing on the Paris station.

The model tends to overestimate ozone concentrations, and underestimate NO₂ and PM₁₀ concentrations. The correlation is good, and the main differences are related to a near-constant bias during the whole day. The most important differences occur during winter, when the boundary layer is less convective and with a lower boundary layer height. The C9 model configuration gives modelled concentrations closer to the measurements than C8 and C20. But the differences between the different model configurations are not so large compared to the differences with the measurements. The most important impact of C9 is for ozone and during summer, reducing the overestimation of the peaks during the afternoon.

6. Discussion and conclusion

This study is dedicated to an evaluation of the impact of different vertical resolutions in modelling near surface concentrations. The 'reference' vertical mesh was defined as the mesh used for the PREVAIR forecast platform using CHIMERE: 8 vertical levels, called C8, ranging from 995 hPa (≈40m above ground level) to 500 hPa (≈ 5000m

agl). Two other meshes were defined: one with the same pressure boundaries (995 to 500 hPa), but with 20 vertical levels (called C20); and one with exactly the same 8 levels, but with an additional level close to the surface, at 999 hPa (≈ 8 m above ground level), called C9.

The quantification of the impact of these different meshes was done by comparing directly the simulations amongst themselves as well as against surface measurements. The domain is the Paris area and the periods studied were chosen to be different in terms of pollution events: the whole month of January 2009 for NO₂ and PM₁₀ pollution, and the whole month of August 2009 for photochemical pollution with ozone, O₃.

Based on monthly averaged maps, it was shown that the differences between C8 and C20 never exceed 3 $\mu\text{g}/\text{m}^3$ during winter and summer and over the whole area. The patterns of differences are mostly related to the main surface emissions sources, here the Paris city, and are slightly spatially shifted of a few kilometers, showing the main wind regimes over the region and the transport of pollutants. In average, a refined mesh (C20) tends to give more NO₂ and PM₁₀ and less ozone than C8. Vertically, C20 is able to capture finer structures, both for gases (O₃) and particles (PM₁₀). These vertical structures have no significant impact on the surface concentrations. The comparison of averaged vertical profiles of vertical diffusivity K_z and pollutants showed that the differences between C8, C20 and C9 are low.

Adding a new point near the surface (C9 simulation) leads to concentrations that are more representative of the height of surface stations measurements. But the differences between C8 and C9 are moderate and cannot fully explain the differences between the model and measurements.

Distributions are calculated to discriminate the concentrations differences as a function of K_z values. Globally, there is no clear signal and the differences seems to be relatively independent of the diffusivity values, except for NO₂ and the lowest K_z values (between 0.1 and 1 m^2/s) and for ozone in case of strong surface diffusivity (i.e $K_z \approx 15 \text{ m}^2/\text{s}$).

Comparisons with surface measurements were performed as time series and diurnal cycles. The main differences between the model and measurements were observed at night, when the vertical diffusivity shows very low values. But this feature is not systematic and, for some cases, low K_z are not the reason for large model to observations differences. The diurnal cycles show an overestimation of surface ozone concentrations and an underestimation of NO₂ and PM₁₀. But for the three vertical meshes, none is able to sufficiently unbiased the model.

Finally, it was shown that different vertical meshes may change of a few $\mu\text{g}/\text{m}^3$ the modelled surface concentrations over a large urbanised area such as Paris. This variability is much smaller than the differences between measurements and predicted values. Changing the vertical mesh may be more realistic in specific cases: high concentrated

ozone layers in the boundary layer, nocturnal concentrations of gas and particles in case of very stable nocturnal boundary layers.

The C8 configuration correspond to the one used in the PREVAIR operational system. We thus showed that the differences when using a refined mesh (C20) or a mesh with a point closest to the surface (C9) are low. This means that results of the C8 operational forecast are not due primarily to the vertical mesh definition.

However, the interest to add a point closest to the surface (C9) has been shown and this correspond to modify the 'urban increment' roughly defined as the difference between the city and background concentrations. This expected change is due to the combined effects of (i) higher deposition rates in remote areas (over vegetative covers) and (ii) higher (or lower for ozone) concentrations over emission areas. The background values seems to be more affected by a change of the whole vertical resolution (simulation C20). This study was conducted over an highly urbanised area, known to be poorly represented in terms of sub-grid scale dynamical processes. While the first layer of C8 and C20 (40m) is more related to the buildings top heights, C9 (8m) is representative of the urban canopy layer and this should be expressed in the vertical mixing representation. Actually, this is not the case in the current CHIMERE version and has certainly to be improved in future model developments.

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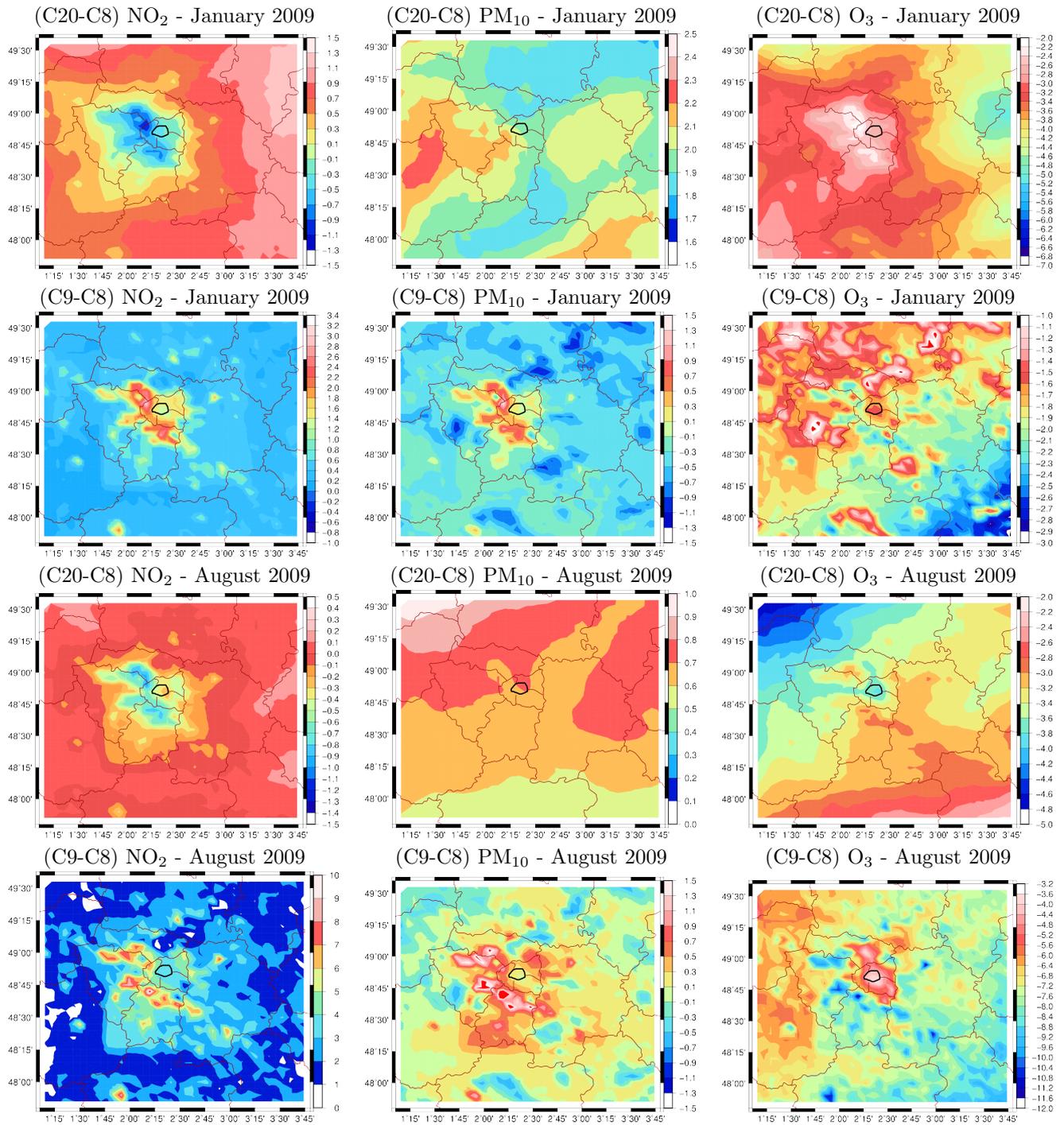


Figure 5: Maps of (C20-C8) and (C9-C8) surface averaged concentrations differences for O_3 , NO_2 and PM_{10} and for January and August 2009. All pollutants concentrations are expressed in $\mu g/m^3$.

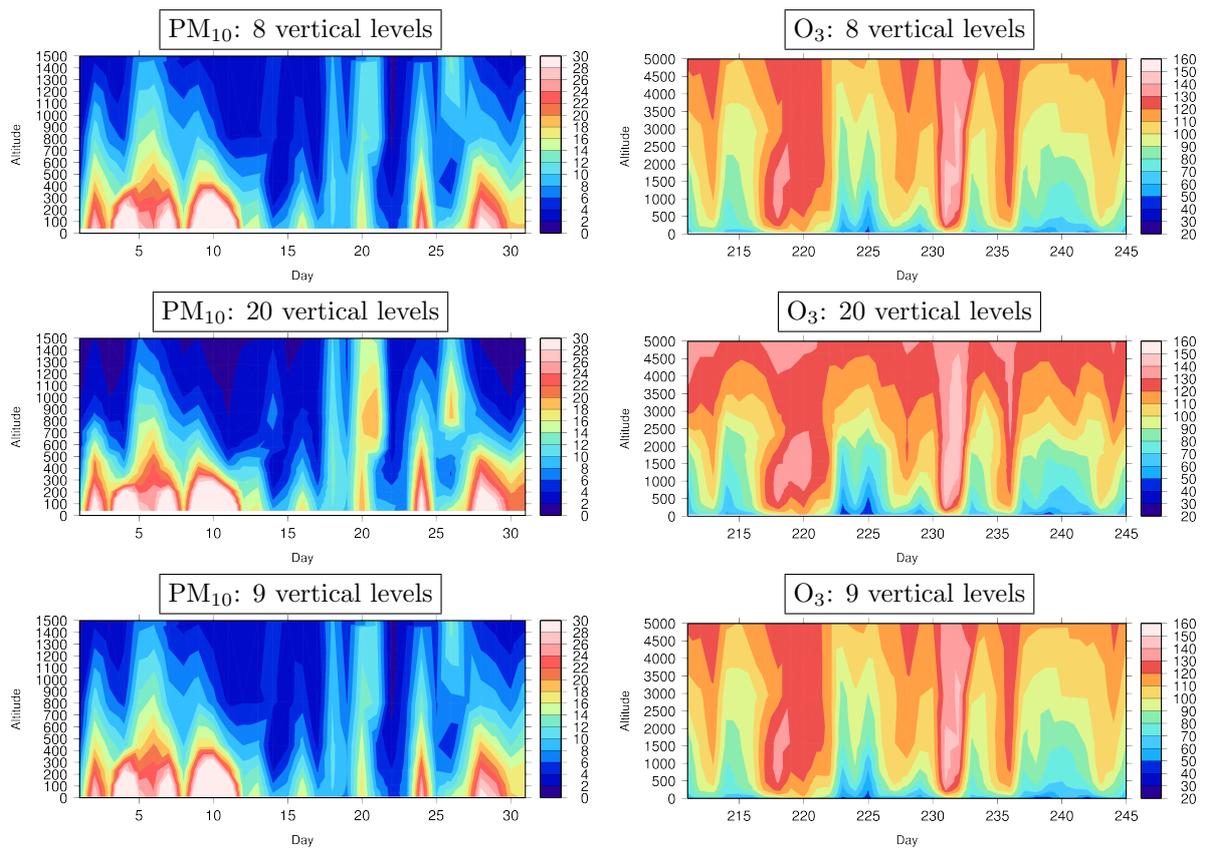


Figure 6: Vertical cross section of [left] PM_{10} concentrations ($\mu g/m^3$) for January 2009 and [right] O_3 concentrations ($\mu g/m^3$) for August 2009.

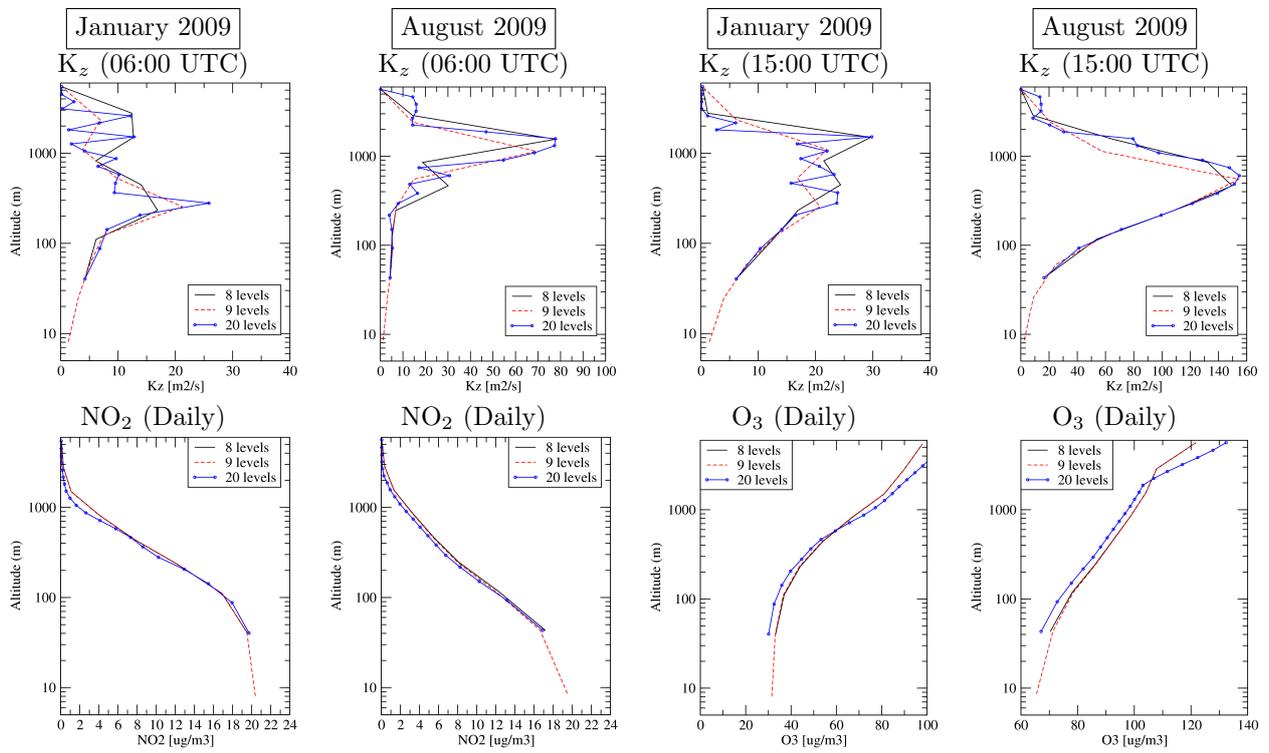


Figure 7: Vertical profiles of K_z , NO_2 and O_3 , for January and August 2009 over Paris. The profiles are averaged for each month of January and August 2009 using only 06:00 and 15:00 UTC for K_z , and all day hours for NO_2 and O_3 .